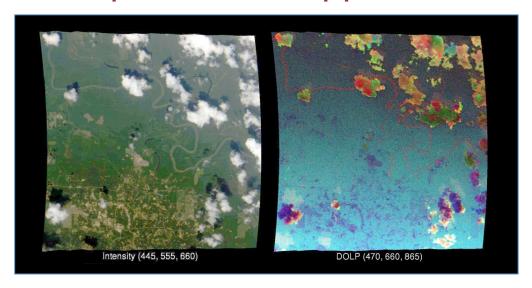
Aerosol and cloud analyses and retrieval algorithm developments in support of MSPI



ACE SWG

Greenbelt, MD

9-11 June 2014

David J. Diner

Jet Propulsion Laboratory, California Institute of Technology

and the MSPI Team



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GroundMSPI and AirMSPI



GroundMSPI is a portable field instrument

2-axis gimbal provides elevation and azimuthal scanning of both the surface and sky

Employed for developing models of surface boundary condition used in aerosol retrievals

Spectral bands: 355, 380, 445, 470*,555, 660*, 865*, 935 nm (*polarimetric)





AirMSPI flies in the nose of NASA's ER-2 aircraft

1-axis gimbal provides multi-angle viewing between ±67°

Being used for developing retrieval algorithms

Aerosol retrieval algorithm development

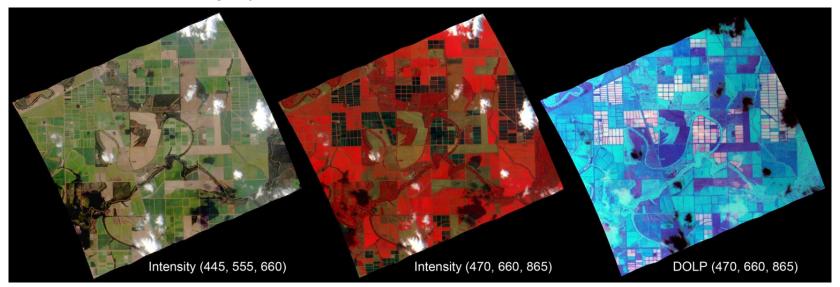
- JPL-developed RT code used as basis of aerosol retrieval algorithm, with support from Oleg Dubovik (Univ. of Lille)
- GRASP code developed by Oleg is being evaluated in parallel

	JPL code (ocean, land)	GRASP (land)
Forward RT calculation method	Markov Chain + Doubling/ Adding	Successive Orders of Scattering
Aerosol size model	Multi-bin, bimodal	Multi-bin, multi-modal*
Particle shape	Spherical	Spherical, spheroidal
Refractive index	Mode dependent	Mode independent
Land surface model	Modified RPV + Fresnel microfacet distribution	RPV + Maignan model
Ocean surface model	Cox-Munk + bio-optical*	Cox-Munk*
Language	Matlab (for development), C++*	Fortran
Speed	Speedup methods required, in study*	Fast

^{*}in development/testing

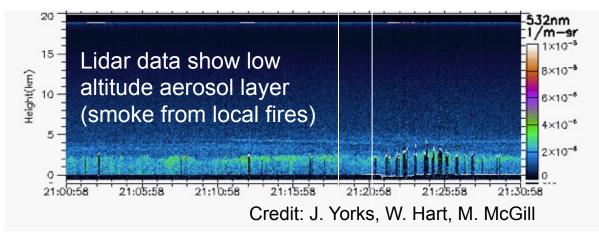
Smoke aerosols near Leland, Mississippi

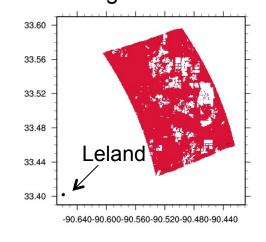
AirMSPI nadir imagery, 9 Sept 2013, 2116 UTC



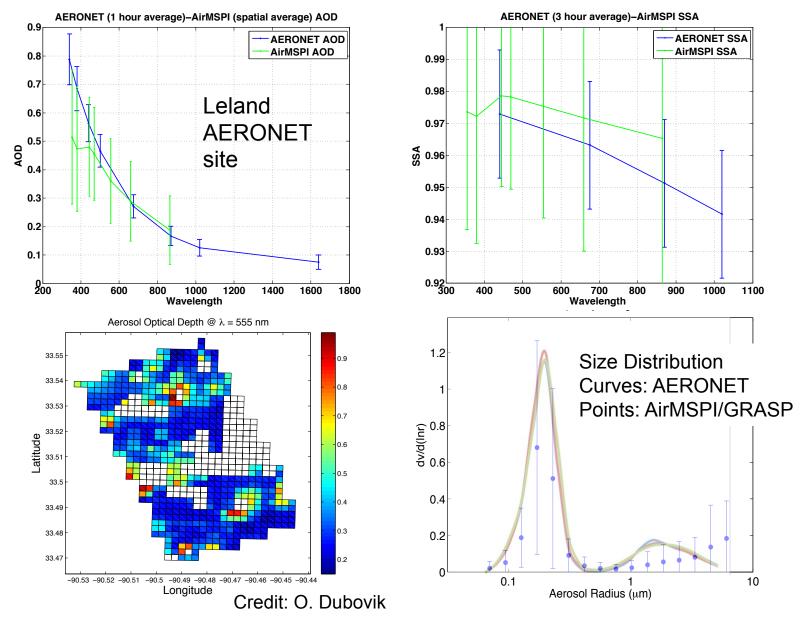
Absorbing Aerosol Index calculated using UV bands A.I. = -100 × $[log_{10}(l_{355}/l_{380})_{meas} - log_{10}(l_{355}/l_{380})_{calc}]$ indicates the presence of absorbing aerosols

CPL backscatter



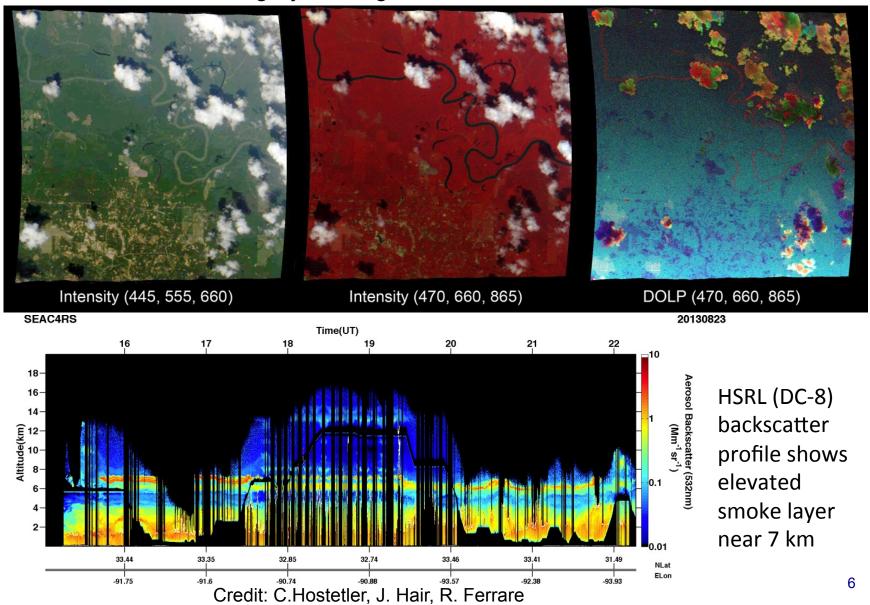


Aerosol retrieval using GRASP

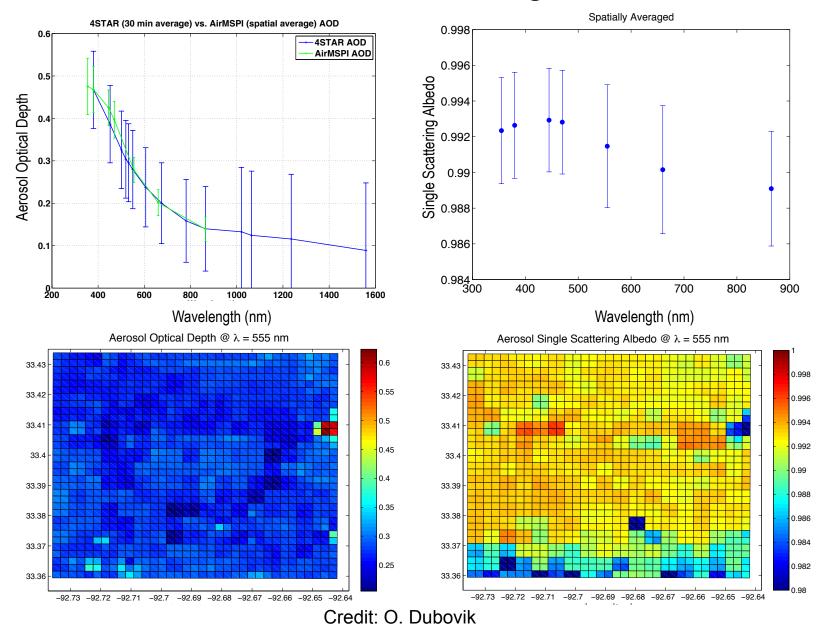


Smoke aerosols in Southern Arkansas

AirMSPI 29.1° aft imagery, 23 Aug 2013, 1636 UTC



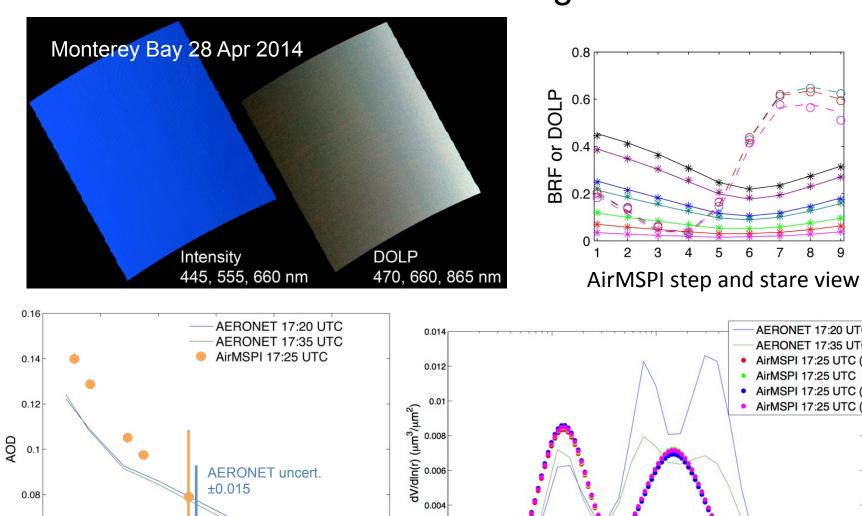
Aerosol retrieval using GRASP



Aerosol retrieval using MarCh

0.002

0.01



±0.015

600

Wavelength (nm)

700

800

900

0.08

0.06

0.04

AirMSPI uncert.

500

±0.030

400

100

5

AERONET 17:20 UTC

AERONET 17:35 UTC

10

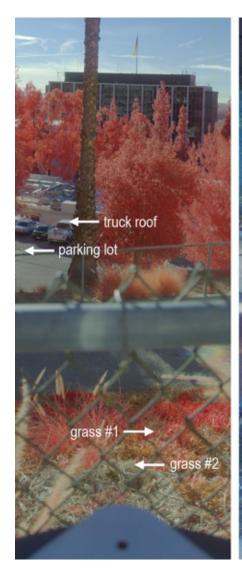
1 Radius (μm)

AirMSPI 17:25 UTC (Patch 1)

AirMSPI 17:25 UTC (Patch 4)

AirMSPI 17:25 UTC (Patch 2) AirMSPI 17:25 UTC (Patch 3)

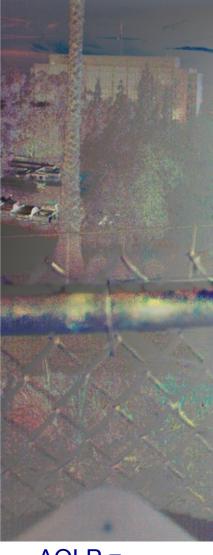
GroundMSPI data analysis



Intensity (I)



DOLP = $(Q^2+U^2)^{1/2}/I$



AOLP = ½ atan (U/Q)

Diner et al. (2012)

470 nm 660 nm 865 nm

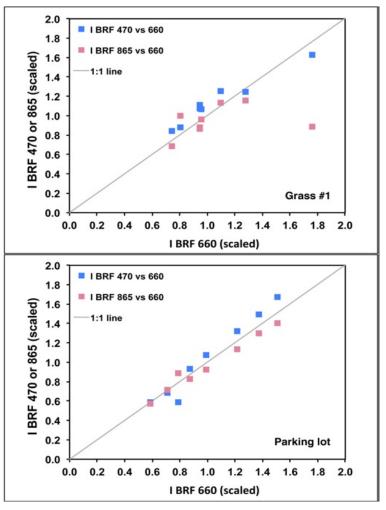
Parametric surface model

Surface BRDF is modeled as a volumetric depolarizing scattering term plus a polarizing term consisting of reflection from an angular distribution of specularly reflecting facets.

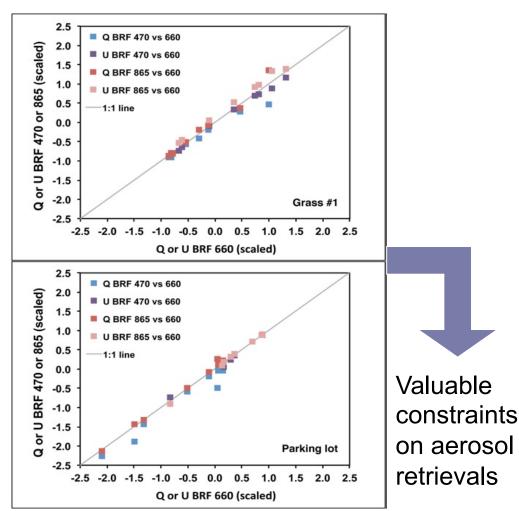
Fresnel facets with angular probability distribution in tilt angle β

Modified Rahman-Pinty-Verstraete

GroundMSPI supports spectral invariance assumption



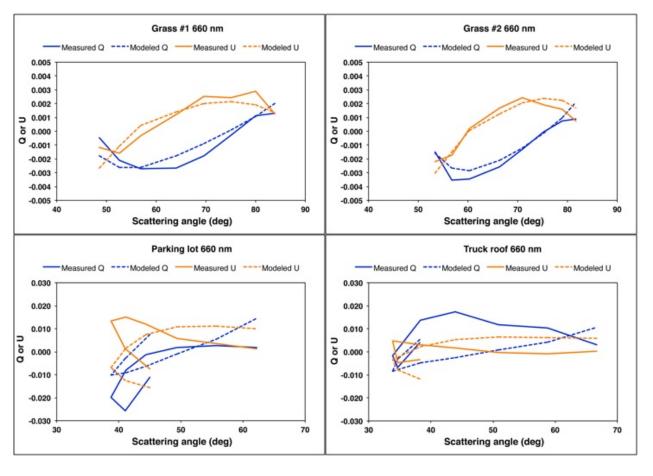
Angular shape of intensity BRF is spectrally invariant (used in MISR retrievals)



Magnitude and angular shape of polarized BRF is spectrally invariant (to be used in MSPI retrievals)

GroundMSPI tests of polarized surface model

Randomly oriented facets reflect sunlight and skylight assuming a single specular (Fresnel) reflection

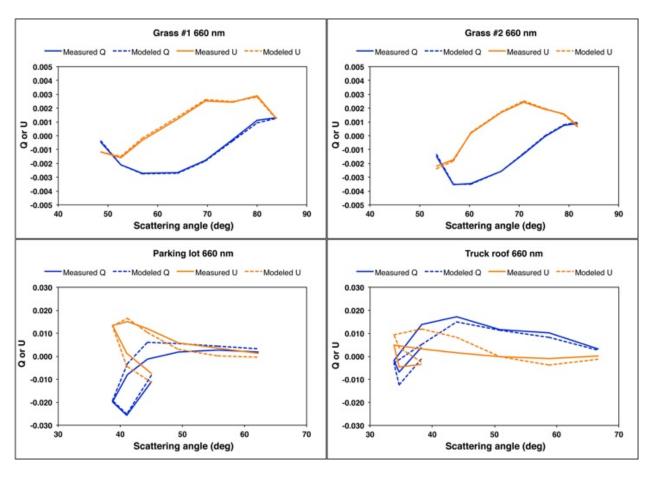


Model predicts correct functional form for natural surfaces

Does not work so well for manmade surfaces

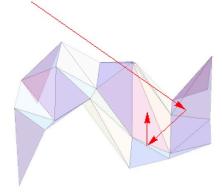
Allowance for multiple scattering

Include empirical spectrally invariant term in surface reflection matrix to allow for rotation of polarization orientation by multiple scattering

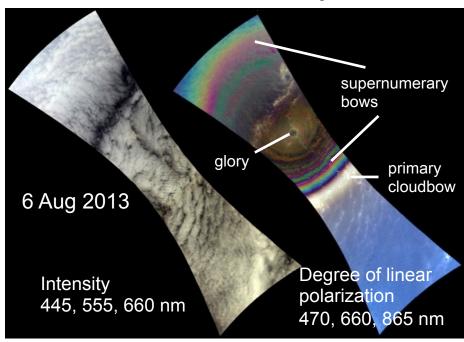


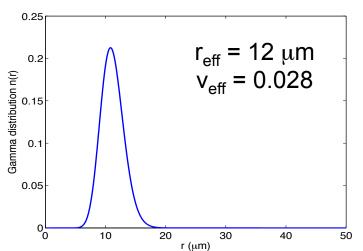
Improves the fit for all surface types

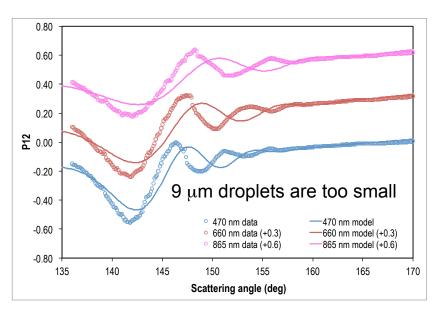
Validity of the model is currently being investigated using polarization ray tracing

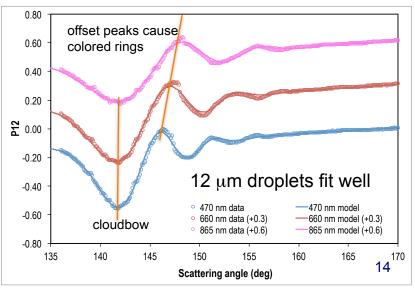


Polarimetric retrieval of cloud-top droplet size distributions

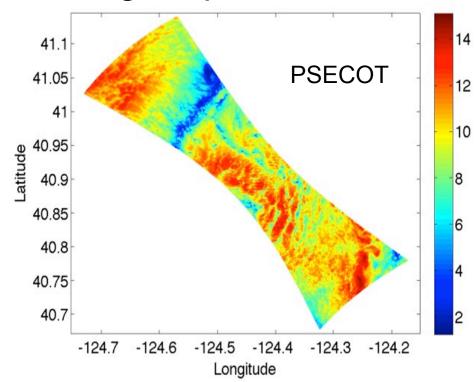








Using drop size to determine bulk cloud properties



 $S_2 = 2^{\text{nd}}$ order structure function: spatial (x,y) and azimuthal (ϕ) average of $|I(x+r_n\cos\phi, y+r_n\sin\phi)-I(x,y)|^2$

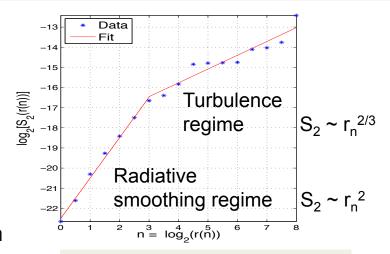
c = value of n at slope discontinuity

$$\eta = p \times 2^c \sim 2 \text{ km} \rightarrow \text{COT} = \langle \text{PSECOT} \rangle_{\text{scale } \eta}$$

Radiative diffusion theory predicts:

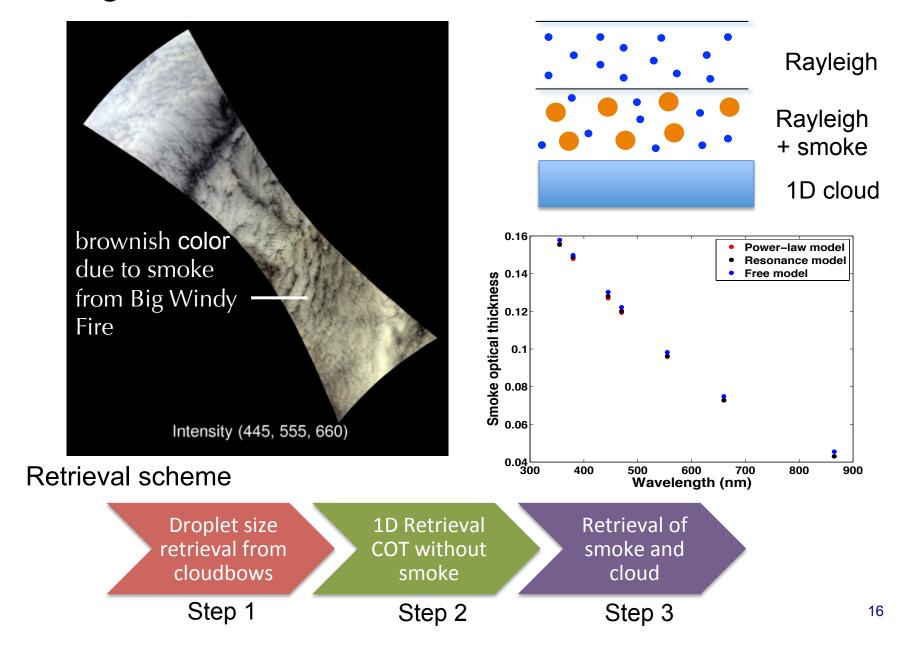
• η is proportional to geometric thickness/COT^{1/2}

- Knowing $r_{\rm eff}$ and $v_{\rm eff}$ from polarized radiance, a look-up table is used to calculate pixel scale effective cloud optical thickness (PSECOT) from the intensity field
- PSECOT is biased by 3D (pixel adjacency) effects, leading to radiative smoothing
- Averaging to determine COT requires determination of the radiative smoothing scale η



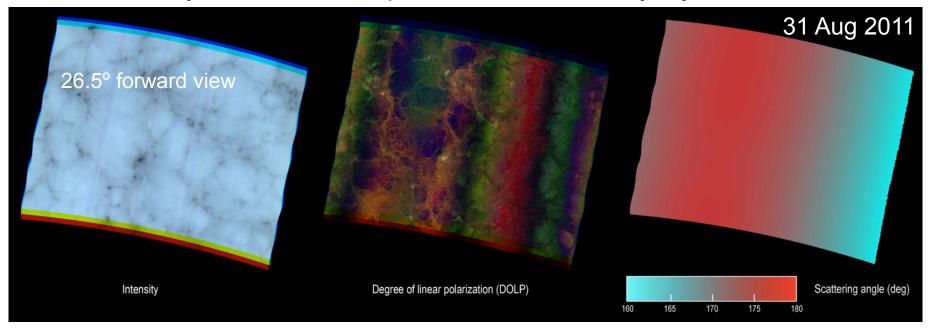
$$r_n = p \times 2^n$$
,
where $p = 25$ m pixel scale

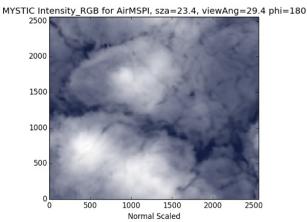
Using cloud model to retrieve aerosol above cloud

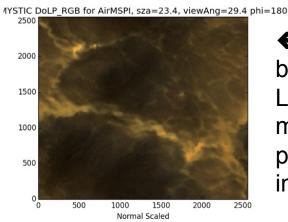


Fine-scale spatial structure in polarized backscatter

Near-backscattering step-and-stare images provide unprecedented view of turbulent dynamics at the tops of marine boundary layer clouds

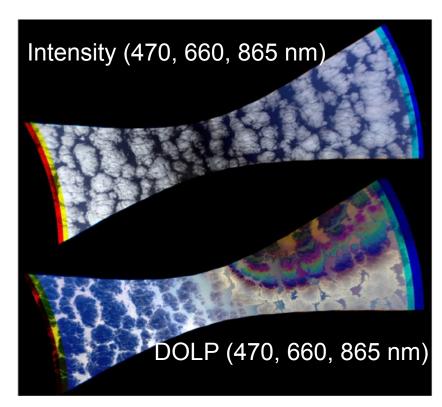






← Synthetic AirMSPI data based on bin-microphysics LES and MYSTIC 3D vRT models reproduces spatial patterns and magnitude of intensity and DOLP

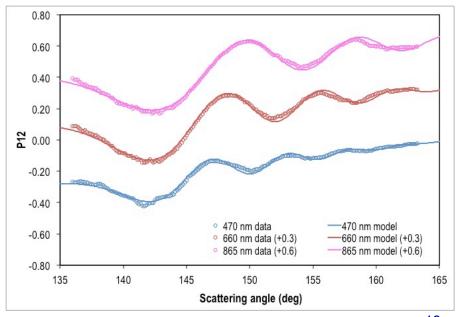
Cloudbow analysis of broken cumulus



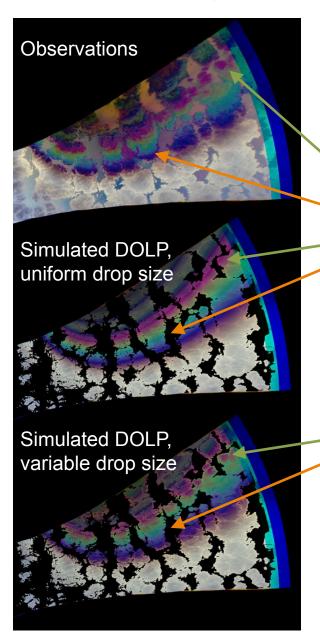
6 February 2013, 2226 UTC - Pacific sweep image

The droplet size retrieval also works for broken clouds

Simple spectral intensity thresholds were used to separate clouds from ocean. Data for the whole scene were fitted with with a distribution having an effective radius of 10 μ m and effective variance of 0.01



Supernumerary bow simulations



Use of a single droplet size does not reproduce the spatially resolved observations, but provides a starting point for a more complete scene model

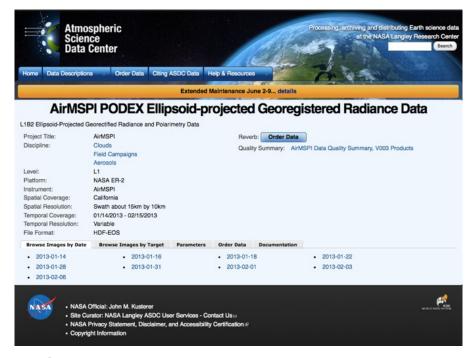
"Scalloped" and "ringed" appearance of supernumerary arcs is not reproduced by assuming constant droplet size of 10 μm

Allowing droplet size to vary from 8 μ m to 10 μ m in proportion to cloud brightness gives a better scene model

This is consistent with smaller sizes at the cloud edges due to evaporation or condensation as the cloud convectively thickens

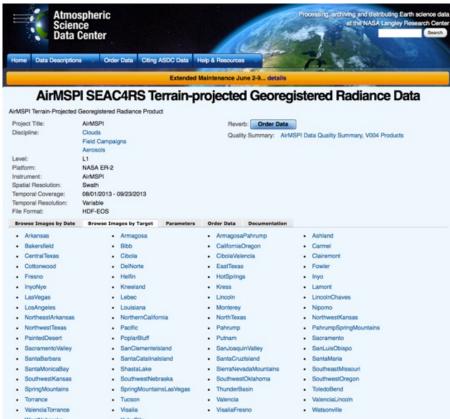
AirMSPI L1 data are available at the LaRC ASDC

https://eosweb.larc.nasa.gov/project/airmspi/airmspi_table



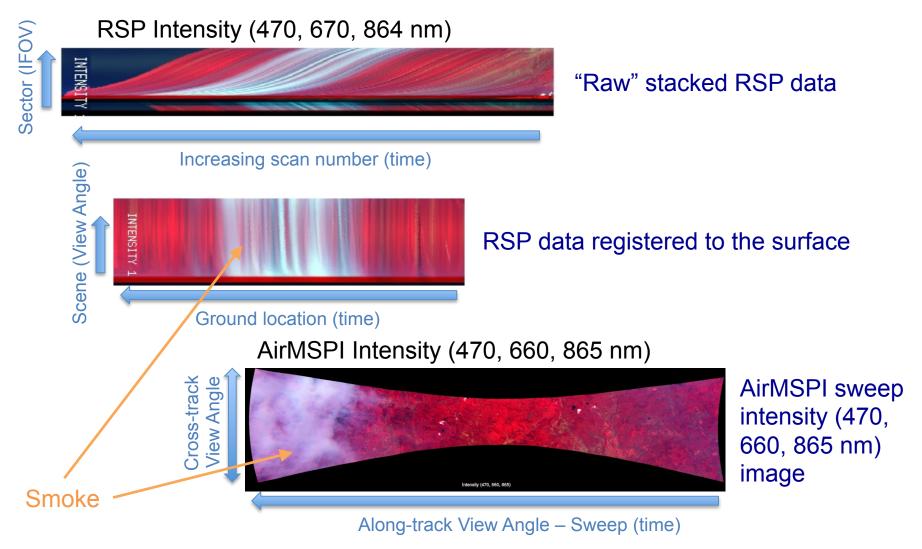
PODEX products (designated v003) were delivered in 2013.

Subsequent PODEX effort has focused on refining radiometric, spectral, geometric, and polarimetric calibration. An updated PODEX delivery will take place in a few months.



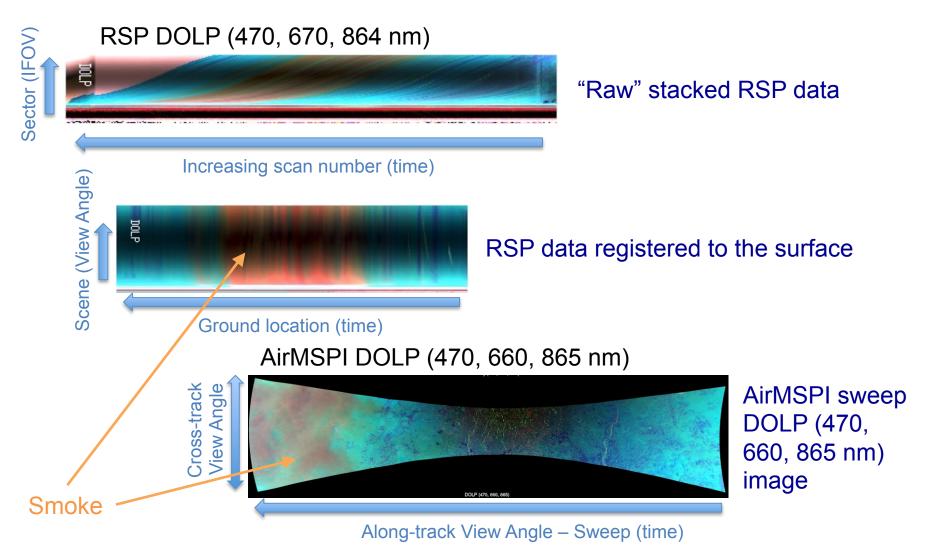
Several calibration improvements were made for SEAC⁴RS. Publicly available SEAC⁴RS products are designated v004.

Example matchup of RSP and AirMSPI intensity



2 August 2013, 2036 UTC – Klamath Mountains (AirMSPI)
2033 UTC (RSP)
RSP data credit: B. Cairns

Example matchup of RSP and AirMSPI DOLP

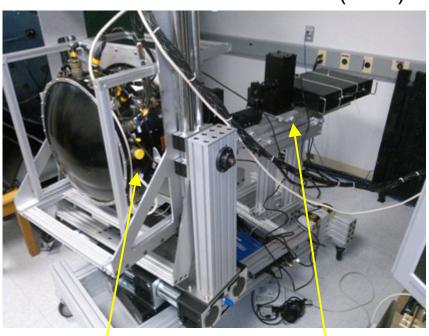


2 August 2013, 2036 UTC – Klamath Mountains (AirMSPI)
2033 UTC (RSP)
RSP data credit: B. Cairns

Laboratory polarimetric cross-calibration of AirMSPI, RSP, PACS?

- Geo-registration of AirMSPI, RSP, and PACS may be a significant source of uncertainty in crosscomparison of flight data
- Propose conducting a laboratory cross-calibration as a controlled experiment
- AirMSPI uses a highly accurate Polarization State Generator (PSG) to polarimetrically calibrate
- Russell Chipman (UofA) has offered to oversee a crosscalibration experiment using the PSG with all 3 instruments, if feasible

Polarization State Generator (PSG)



AirMSPI mounted on automated stage

PSG mounted on fixed rail

Acknowledgments

AirMSPI engineering and operations

Brian Rheingans, Sven Geier, Sebastian Val, Steve Adams (JPL), Karlton Crabtree (Univ. of AZ)

AirMSPI science

Michael Garay, Olga Kalashnikova, Anthony Davis, Michael Tosca (JPL), Larry Di Girolamo (Univ. of IL), Ralph Kahn (GSFC), Roger Marchand (Univ. of WA)

AirMSPI calibration

Carol Bruegge, Felix Seidel (JPL), Brian Daugherty, Russell Chipman (Univ. of AZ), Ab Davis (Univ. of TX)

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Flight ops

Stu Broce, Dean Neeley, Tim Williams, Denis Steele, Tim Moes, Chris Miller (AFRC)

Sponsors

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